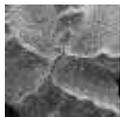


Upper Silurian and Lower Devonian conodonts from the Monte Cocco II Section (Carnic Alps, Italy)

MARIA G. CORRIGA & CARLO CORRADINI



The upper Silurian-Lower Devonian Monte Cocco II section, located in the eastern part of the Carnic Alps, yielded abundant conodonts from five conodont biozones (*snajdri*, *crispa*, *eosteinhornensis* s.l., *detortus* and *woschmidti*). The *eosteinhornensis* s.l. interval Zone is introduced here to replace the “*Oz. remscheidensis* i.Z.” of Corradini & Serpagli (1999). The Silurian/Devonian boundary is located in the upper part of the section; problems related to the accurate location of the Silurian/Devonian boundary on the basis of conodonts are discussed. A new taxon of the genus *Pelekysgnathus* is described, but left in open nomenclature. A few taxonomic notes on other selected taxa are also presented.

• Key words: Silurian/Devonian boundary, conodonts, biostratigraphy, taxonomy, Carnic Alps.

CORRIGA, M.G. & CORRADINI, C. 2009. Upper Silurian and Lower Devonian conodonts from the Monte Cocco II Section (Carnic Alps, Italy). *Bulletin of Geosciences* 84(1), 155–168 (6 figures, 1 table). Czech Geological Survey, Prague. ISSN 1214-1119. Manuscript received November 24, 2008; accepted in revised form March 10, 2009; published online March 23, 2009; issued March 31, 2009.

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This paper is a contribution to a cooperative research project on the Silurian of the Carnic Alps arranged between several Italian universities (Cagliari, Modena and Reggio Emilia, Pescara) and institutions (Museo Friulano di Storia Naturale). Research includes geological mapping, palaeontology, stratigraphy and depositional history of Silurian rocks, with special regard to carbonatic facies, in several sectors of the Carnic Chain. The ultimate goal of the project is to provide a reconstruction of the paleoenvironmental and paleogeographical evolution of the Carnic Basin during Silurian times.

Silurian and Lower Devonian sediments in the Carnic Alps are represented either by cephalopod limestone (“*Orthoceras* limestone”) or black shales. They crop out discontinuously in many localities in the region between Wolayer Lake and the Uqua Valley. The central and western sectors of the chain (Wolayer Lake, Passo di Monte Croce Carnico/Plöckenpaß and Oberbuchach areas) have been investigated and described in great detail by several authors, while only a few papers deal with Silurian topics from other areas of the Carnic Alps.

Monte Cocco is located in the eastern part of the Carnic Alps. Here Silurian and Lower Devonian sediments are mainly represented by “*Orthoceras* limestones”, but locating sections spanning the Silurian/Devonian boundary is difficult, because of either tectonic disturbance or the Quaternary cover. The “Monte Cocco II” section is the only

section in the area where the Silurian/Devonian boundary is exposed and yields a quite rich conodont fauna.

Revision of conodont stratigraphy across the S/D boundary has been provided recently by Carls *et al.* (2007). In addition to defining some new taxa, the authors highlighted and discussed in detail problems with the various biozonation schemes proposed for the Přídolí time. Although their arguments are valid, unfortunately they do not provide any new scheme to substitute for those criticized. Also, the recognition of the Silurian/Devonian boundary on the basis of conodonts is still a problem; the boundary is defined by the first occurrence of the graptolite species *Mongraptus uniformis* with no conodont taxa appearing at exactly the same time

Previous work in the area

The Monte Cocco area has long been known to geologists because of the abundance of fossils (Frech 1888) and the presence of iron and manganese which led to mining activity (Feruglio 1970, Zucchini 1998). However, geologic and stratigraphic studies of this area are quite old, with the exceptions of Herzog (1988), which has a structural approach only and Corradini *et al.* (2003), which is a preliminary report on the Silurian Cephalopod limestones. It should be pointed out that this locality is referred to by its Austrian name, “Kokberg”, in old papers and by German speaking

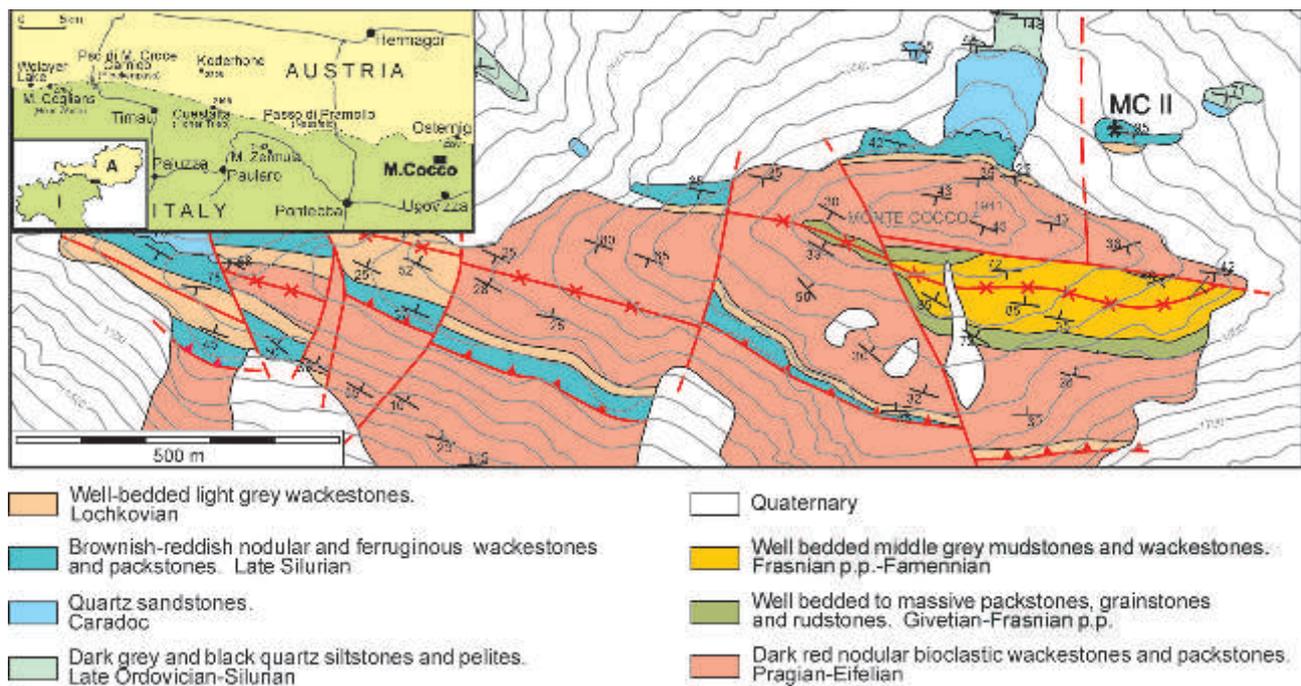


Figure 1. Geological map of the Monte Cocco area (modified after Corradini *et al.* 2003), with the location of the Monte Cocco II (MC II) section.

scientists; Monte Cocco was annexed to Italy only at the end of World War I.

The Silurian rocks of Monte Cocco are highly fossiliferous, and many papers deal with the taxonomy of selected groups, especially the nautiloid cephalopods (for a list of papers see Corradini *et al.* 2003). A few bivalves have been listed by Kříž (1999) in his monograph on Silurian bivalves from the Carnic Alps. Manzoni (1965) and Herzog (1988) looked at the occurrence of conodonts and listed some taxa of Wenlock and Ludlow age, and Corradini *et al.* (2003) published a preliminary paper on conodont stratigraphy of Ludlow-Přídolí age, in which some macrofossils were also illustrated.

Geological setting

Sediments of Late Ordovician to Late Devonian age are exposed in the Monte Cocco area (Fig. 1). The basal part of the succession consists of dark grey to greenish pelites and siltstones, grading to light grey sandstones (Uqua Fm.) of Katian-Hirnantian age. Locally, a massive 10 m thick crinoidal limestone of presumably Hirnantian age occurs.

The Silurian rocks are relatively widespread and mainly represented by a cephalopod limestone in Wolayer facies (Corradini *et al.* 2003), which disconformably overlies the Ordovician sediments. The lower part of the Silurian sequences (up to the uppermost Gorstian) is represented by dark brown to black wackestone to packstone (Kok Fm.), cropping out mainly on the western flank of

Monte Cocco. The upper part of the succession consists of about twenty meters of wackestones and packstones (“Alticola Lms”) that are well stratified in beds 10–50 cm thick. The brownish-reddish colour of these beds frequently weathers to dark red due to the abundance of iron minerals present, and progressively grades to light grey across the Silurian/Devonian boundary. Fossils are common, mainly nautiloid cephalopods, trilobites and bivalves (Corradini *et al.* 2003). The inference of pelagic and hemipelagic deposition for these beds is supported by the lack of structures documenting traction currents.

The “*Orthoceras* limestones” are conformably overlain by about 10 m of well-bedded light grey Lochkovian cephalopod wackestones and packstones (Rauchkofel Fm. and Boden Lms), which represent a marked shallowing of the sea level. The depositional processes are still pelagic and hemipelagic, but a few tempestites reflecting major storm events are documented.

The succession continues with dark red nodular limestones (Findenig Fm.) of Pragian to Givetian age and bioclastic grainstones and rudstones (Eifelian-Givetian). The youngest pre-Variscan sediments in the Monte Cocco area are represented by pelagic limestones (“*Clymenia* Lms”) of Upper Devonian age.

The Monte Cocco II (MC II) section

The Monte Cocco II section is located on the northern flank of Monte Cocco at GPS coordinates N 46°33′ 03”, E

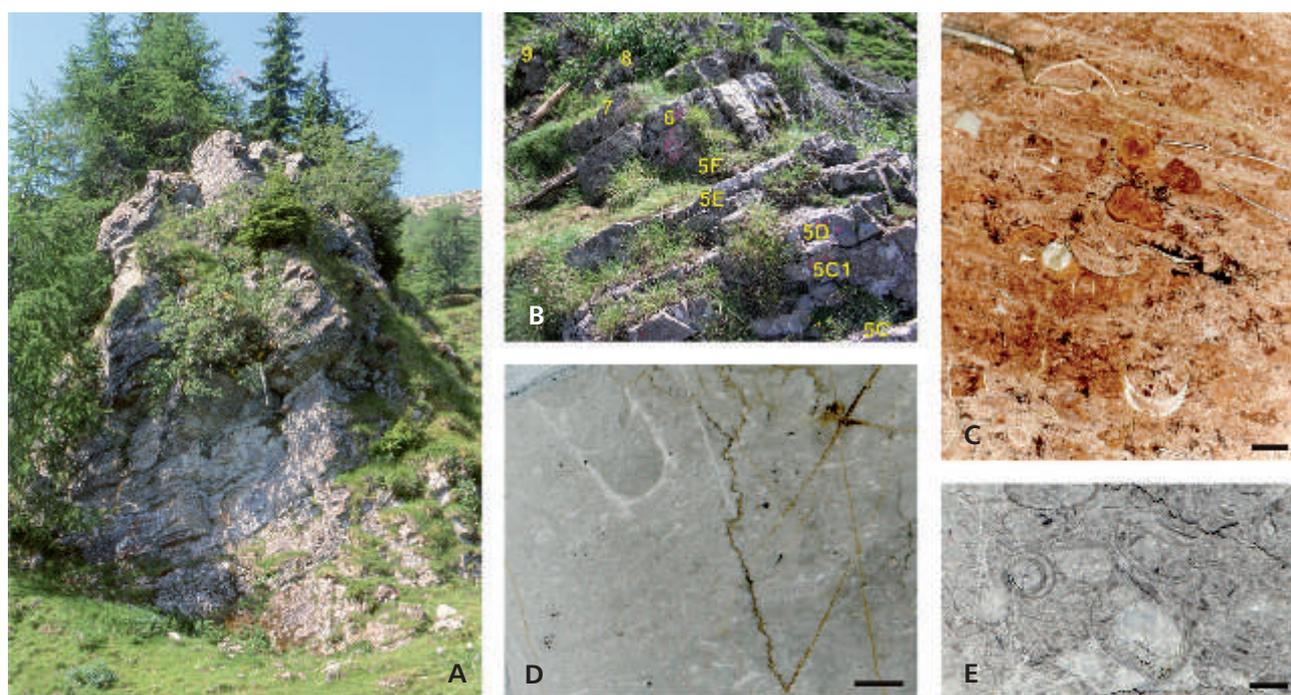


Figure 2. A – panoramic view of the Monte Cocco II section. • B – the upper part of the Monte Cocco II section: the Silurian-Devonian boundary is likely between beds 5D and 5E. • C – bioclastic wackestone, sample, MC II 2, *crispa* Zone; scale bar = 2 mm. • D – bioclastic grainstone with a large orthoceratid shell, sample MC II 6, *woschmidti* Zone; scale bar = 2 mm. • E – bioclastic wackestone, sample MC II 5, *detortus* Zone; scale bar = 2 mm.

13°26' 47" (Fig. 1). The section consists of about 17 metres of *Orthoceras* limestone (Figs 2, 3) that are divided into three lithostratigraphic units: "Alticola Limestone", Rauchkofel Formation and Boden Limestone.

The lower part of the section exposes about 14 m of "Alticola Limestone": a grey-pale brownish cephalopod wackestone, which in some intervals turns dark red due to weathering of abundant iron minerals. The upper part of the unit is more greyish in colour and progressively grades to the grey packstone of the neritic Rauchkofel Fm. (1.5 m). The upper part of the section is represented by about 1.5 m of well-bedded, pale-grey-whitish limestone of the Boden Limestone. However, it should be pointed out that it is difficult to precisely locate unit boundaries because they are gradational and also because of weathering.

Nautiloid cephalopods, trilobites and bivalves are abundant throughout the section, whereas crinoids, gastropods and brachiopods are rare. The dominant microfacies is cephalopod wackestone and packstone (Fig. 2C–E). Thin-section analysis revealed echinoderms, gastropods, and cephalopods; small shells (ostracods, bivalves and brachiopods) were also present.

Conodont fauna

Twenty-six samples were processed with the conventional formic acid technique. More than 65 kg of limestone yielded

about 1700 conodont elements (Table 1). The state of preservation is generally quite good, even if a few elements are broken or slightly deformed. The abundance varies from 136 elements/kg (sample MC II 1A) to 2 elements/kg (sample MC II 8), with an average of about 25 elements/kg. In general, the lower part of the section has a much greater abundance of conodont elements, while the upper part (Boden Lms) has low abundance. Conodont colour is dark brown, corresponding to a Colour Alteration Index of 3.5–4.

Twenty-seven taxa belonging to fourteen genera were identified: *Belodella*, *Coryssognathus*, *Dapsilodus*, *Dvorakia*, *Icriodus*, *Lanea*, *Oulodus*, *Ozarkodina*, *Panderodus*, *Pedavis*, *Pelekysgnathus*, *Pseudooneotodus*, *Wurmiella* and *Zieglerodina*.

Ozarkodinids are always very abundant throughout the section, representing approximately one half of the entire fauna. Coniforms (mainly *Dapsilodus* and *Panderodus*) are only dominant in the lower part of the section, up to sample MC II 1A.

Biostratigraphy

The biozonation schemes followed in this paper are those proposed by Corradini & Serpagli (1999) for the Silurian and by Carls & Weddige (1996) for the Lower Devonian. However, it is necessary to point out that the Přídolí zonation of Corradini *et al.* (1998), Corradini & Serpagli (1999)

Table 1. Distribution of conodonts in the Monte Cocco II section

		0	1	1A	1B	1C	2	3	3A	3C	3D	4	4A	4B	5	5A	5B	5C	5C1	5D	5E	5F	6	7	8	9	total
<i>Belodella resima</i>		2								2		5	6	1	2		3					6	13				40
<i>Belodella anomalis</i>											1		1						1								3
<i>Corysognathus dubius</i>	Pc														1												1
	M											3			2												5
	Sa/Sb						1					2	3														6
	Sc						1	1				1															3
	conif											3			2												5
<i>Dapsilodus obliquicostatus</i>		42																									42
<i>Dvorakia</i> sp.																								1			1
<i>Icriodus w. hesperius</i>	Pa																					1	3				4
<i>Icriodus w. woschmiditi</i>	Pa																					1					1
<i>Icriodus</i> sp.	Pa																						1				1
<i>Lanea omoalpha</i>	Pa																						2				2
<i>Oulodus elegans elegans</i>	Pb									1				1	1												3
	M					1									1						1						3
	Sa								1	1					1												3
	Sb					1			1						2				1								5
	Sc					1									1												2
<i>Oulodus el. detortus elegans detortus</i>	Pb										1		1														2
	M													1													1
	Sa													1													1
	Sb												1														1
	Sc										1			1													2
<i>Oulodus siluricus</i>	M	1																									1
	Sa	1	1															1									3
	Sb	2																	1								3
	Sc	1			1													1									3
<i>Ozarkodina crispa</i>	Pa		6																								6
<i>Ozarkodina confluens</i>	Pa	7											1											3			11
	Pb												2														2
	M												2														2
	Sa	1																									1
	Sc	2											1														3

		0	1	1A	1B	1C	2	3	3A	3C	3D	4	4A	4B	5	5A	5B	5C	5C1	5D	5E	5F	6	7	8	9	Total	
<i>Ozarkodina cf. confluens</i>	Pa												1															1
<i>Ozarkodina eosteinhornensis</i> s.s.	Pa										15	12																27
<i>Ozarkodina eosteinhornensis</i> s.l.	Pa					2					3				15					1	1		2					24
	Pb					2					5	2			13								2					24
	Sa					1						2																3
<i>Ozarkodina planilingua</i>	Pa														1							2						3
<i>Ozarkodina</i> sp.	Pa																3											3
<i>Panderodus recurvatus</i>		2	4	2																								8
<i>Panderodus unicastatus</i>		65	163	266	1		1						13	6				7	1				1					524
<i>Pedavis biexoramus</i>	Pa																					3	3					6
<i>Pelekysgnathus</i> sp. A	Pa		1																									1
<i>Pseudooneotodus beckmanni</i>		1	6		3		8	1		1		5	26	3	5	1	1	2				5	38	4				110
<i>Ps. bicornis contiguus</i>					2																							2
<i>Wurmiella excavata</i>	Pa	15	37	1	5	5	14	4	1	26	7	2	9	2	2		1	7	3									141
	Pb	6	12	2	2	1	5	3		14	2	5	6	2	7		1	2										70
	M	4	8	3	1	1	5	2		21	1	3	4	1	2		1	3										60
	Sa	9	7		3		11	3	1	28	1	2	4	1	2		1	6										79
	Sb	9	16		3		15	3	1	26	3	7	2	2	5		1	5	1									99
	Sc	10	24	3	4	2	21	4		26	5		8	3	6	1	1	9	2									129
<i>Zieglerodina remscheidensis</i>	Pa														6						1	3	2	4	1	4		21
	Pb														8						2		1	1	1			13
	M														1									2				3
	Sa																					1		1				2
	Sc														3						2		1		1			7
Indetermined coniforms										1					15	1	1	2				3						23
Indetermined ramiforms		1	2		3					1	3		3	3	4	4	1											25
Fragments		4	23	2	2	1	13	1	2	5	3		1	6	16	1	4	8	9	2	5	4	3	3	2			120
Total		185	310	279	30	18	95	22	7	153	48	57	92	34	123	8	23	54	18	6	12	29	72	16	4	5		1700
kg rock		1.6	5.1	2.05	2.36	1.55	5.1	4.5	1.25	3.65	1.66	4.15	3.95	1.5	4.7	1.44	2.41	4.35	2.76	1.23	0.93	1.3	3.0	1.83	1.6	1.8		65.77
conodonts/kg		115.6	60.8	136	12.7	11.6	18.6	4.9	5.6	41.9	28.9	13.7	23.3	22.7	26.2	5.6	9.5	12.4	6.5	4.9	12.9	22.3	24	8.2	2.5	2.8		25.8

should be updated, as already suggested by Corrigan *et al.* (2009). The authors named an “*Oz. rem. remscheidensis* interval Zone” in the lower Přídolí, defining an interval between the last occurrence of *Ozarkodina crispera* and the first occurrence of *Oulodus elegans detortus*. However, recent taxonomic revisions of late Silurian Ozarkodinids by Murphy *et al.* (2004) and Carls *et al.* (2007) state that *Zieglerodina remscheidensis* (Ziegler) is restricted to the Lower Devonian and does not occur in the lower part of Přídolí. Therefore, it is not appropriate to name a zone by an absent taxon. Awaiting a careful taxonomic revision of Přídolian Ozarkodinids, we choose to rename the “*Oz. rem. remscheidensis* interval Zone” of Corradini & Serpagli (1998, 1999) after the more common taxon: *Ozarkodina eosteinhornensis* s.l. It should be pointed out that this name variation does not change the biostratigraphic meaning of the zone, since the definition of the boundaries remains the same.

The conodont association allows recognition of five conodont zones in the Monte Cocco II Section: *snajdri*, *crispera*, *eosteinhornensis* s.l. and *detortus* Zones in the Silurian and the *woschmidti* Zone in the Devonian.

The snajdri Interval Zone. – No markers are present in sample MC II 0. However, the occurrence of *Ozarkodina crispera* (marker of the *crispera* Zone) in sample MC II 1 allows us to state that the strata below belong to the immediately older *snajdri* Zone. According to Corradini & Serpagli (1999) the upper limit of the *snajdri* interval Zone is defined by the first occurrence of *Oz. crispera*. *Panderodus unicosatus* and *Dapsilodus obliquicosatus* are abundant in this interval; a similar abundance of these taxa also occur in coeval levels in the Cellon section, and may be related to regional ecological factors.

The crispera Zone. – The lower boundary of the *crispera* Zone is defined by the first occurrence of the marker *Oz. crispera*, and the upper boundary by the last occurrence of the same taxon. It has been found only in sample MC II 1 and therefore only this level belongs to the *crispera* Zone. *Panderodus unicosatus* is dominant, and *Wurmiella excavata* is also very abundant.

The eosteinhornensis s.l. interval Zone. – The *eosteinhornensis* s.l. Zone is distinguished from sample 1A to sample 3C in that the lower boundary is defined by the last occurrence of *Oz. crispera*, and the upper boundary is defined by the first occurrence of *Oulodus el. detortus*. *Wurmiella excavata* is largely dominant in all the levels of this interval,

with the exception of sample MC II 1A, where *Panderodus unicosatus* represents more than 90% of the association.

The detortus Zone. – The *detortus* Zone occurs in samples 3D to 5E. The lower boundary is defined by the first occurrence of *Oulodus elegans detortus*, and the upper boundary by the entry of Icriodids (either *I. woschmidti* or *I. hesperius*). The marker has not been found in the upper part of the Zone, since it has its last occurrence in sample MC II 4B.

For a long time after Jeppsson’s (1988) paper on the Silurian/Devonian GSSP, the Klouk section, *Oul. el. detortus* was considered the uppermost Přídolí marker, just below the Silurian/Devonian boundary. However, it has been recently demonstrated that this taxon already occurs by around the middle of the Přídolí in several geographic areas, such as Sardinia (Gouwy & Corradini 2006), the Czech Republic (Požáry section; Carls *et al.* 2007) and Frankenwald (Carls *et al.* 2007). This is also the case in the Monte Cocco II section; therefore this zone is also definitely longer in the Carnic Alps.

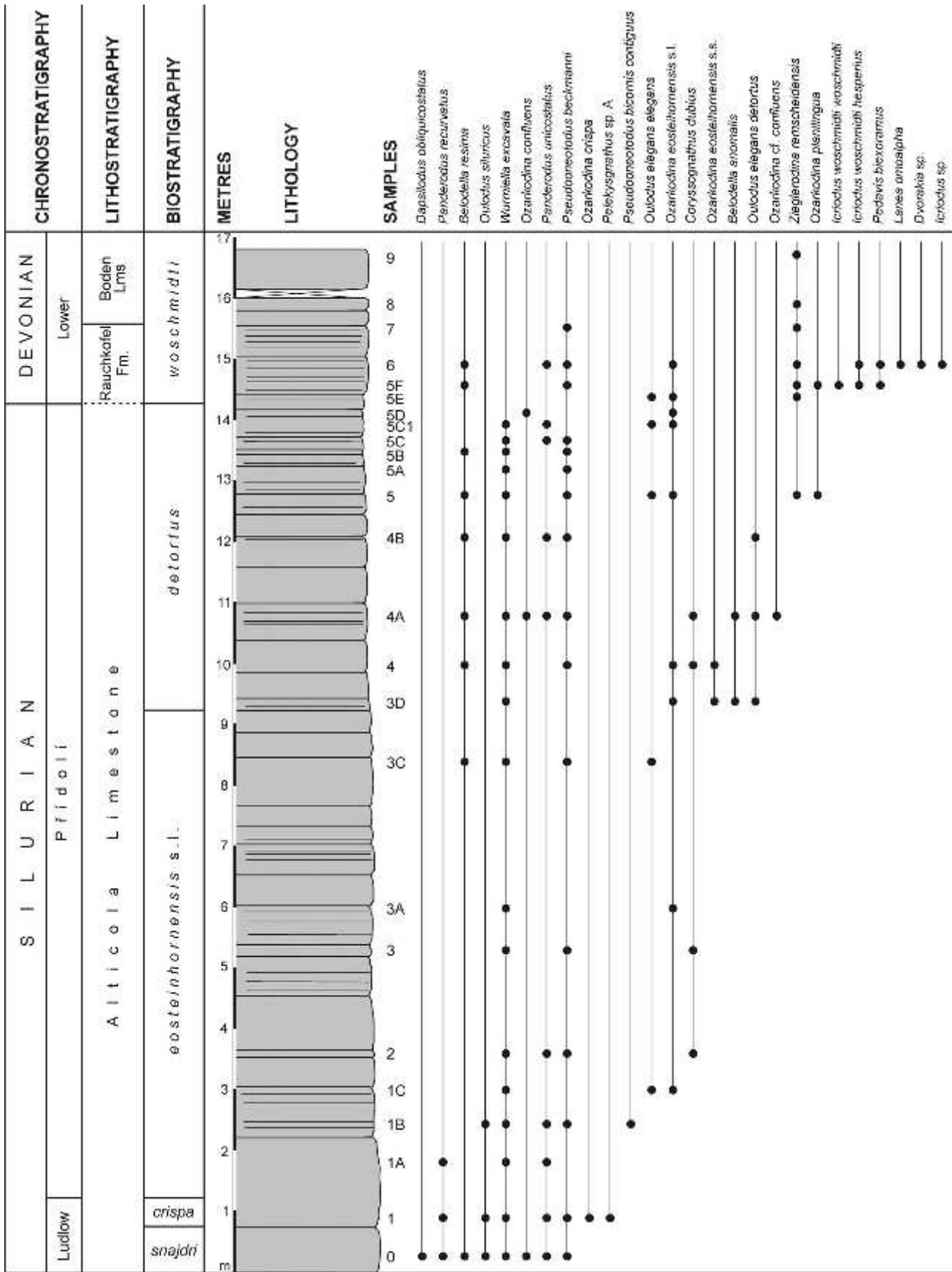
The woschmidti Zone. – The joint first occurrence of *I. woschmidti* and *I. hesperius* in sample 5F allows us to refer this beds and the upper part of the section to the *woschmidti* Zone. However, the entry of *Pe. biexoramus* in the same sample suggests that this level is in the high part of the zone. For discussion on the Silurian/Devonian boundary see below.

Comments on the Silurian/Devonian boundary

Recognizing the Silurian/Devonian boundary on the basis of conodonts has always been a problem because the boundary is defined by the first occurrence of the graptolite *Monograptus uniformis*, and no known conodont taxon has a simultaneous first occurrence.

Icriodus woschmidti is the marker of the first conodont zone of the Lochkovian and its presence is often considered as a proof of Devonian age. However, according to Jeppsson (1988), in many sections *I. woschmidti* occurs slightly before *M. uniformis*, but this fact may be due to facies control: graptolites are normally collected from shales and conodonts from limestones. Furthermore Icriodids are typical of inshore shallow water deposits (Bultynck 2003), but are rare in other marine environments. As a result *I. woschmidti* is very rare, if not absent, in many sections. It should be pointed out that possible problems in the discrimination of *I. woschmidti* from *I. hesperius* may occur

Figure 3. Stratigraphic log of the Monte Cocco II section, with indication of stratigraphic units and occurrence of main conodont taxa. Numbers on the right of the column indicate conodont samples.



and in several localities the latter species has an older first occurrence (M. Murphy, pers. comm.). Therefore the entry of one or the other *Icriodus* taxon may be used for recognizing a basal Devonian age.

Jeppsson (1988, fig. 1) draws the S/D boundary more or less coincident with the top of the *L. e. detorta* Zone (here *detortus* Zone). In some areas *Oulodus elegans detortus* has a longer range within the Přídolí (Gouwy & Corradini 2006, Carls *et al.* 2007), but *Oul. el. detortus* always becomes extinct just below the base of the Devonian and before the entry of *Icriodus woschmidti* and/or *I. hesperius*.

The Silurian/Devonian boundary, therefore, should be located between the last occurrence of *Oulodus elegans detortus* and the first occurrence of *Icriodus woschmidti* and/or *I. hesperius*. In the Monte Cocco II section *Oul. el. detortus* has its last occurrence in sample MC II 4B, whereas *Icr. hesperius* occurs 2.5 metres above in sample MC II 5F together with *Pedavis biexoramus*. The latter species has been reported somewhat higher within the *uniformis* graptolite Zone, after the last occurrence of *I. hesperius* (Schönlaub 1980) or slightly higher (*eurekaensis* Zone of Murphy & Matti 1982); therefore that level may not correspond to the base of the *woschmidti* Zone.

Oz. confluens (= *Oz. typica sensu* Murphy *et al.*, 2004) does not reach the Devonian (M. Murphy, pers. comm.), and in the Monte Cocco II section is discontinuously present up to sample MC II 5D. *Zieglerodina remscheidensis* appears in sample MC II 5 and then is present continuously from MC II 5E and above. Murphy *et al.* (2004) introduced the genus *Zieglerodina* and revised the type species, stating that *Z. remscheidensis* has a restricted range in the early Lochkovian, but gave no key for determination of many other morphotypes strictly related to this species and previously identified as "*Oz. remscheidensis*". Our material from sample MC II 5 is represented by poorly preserved small elements, whereas specimens from sample MC II 5E show typical features of *Z. remscheidensis sensu* Murphy *et al.* (2004). Thus, it is possible that in that restricted taxonomic interpretation and after a complete revision of species belonging to genus *Zieglerodina*, the specimens from sample MC II 5 represent some early form of *Z. remscheidensis*, whereas late forms of the species occur near the base of the Devonian. Unfortunately, our material is not preserved well enough to make any decisive conclusion.

On the basis of our findings, the Silurian/Devonian boundary can be tentatively placed in the Monte Cocco II section between sample MC II 5D and sample MC II 5E.

Systematic palaeontology

Systematic notes are restricted to necessary taxonomic or biostratigraphic remarks. For suprageneric classification, the scheme proposed by Sweet (1988) is followed.

The studied material is stored in the Earth Science Department of the University of Cagliari; figured specimens are housed in the Palaeontological Museum of the University of Modena and Reggio Emilia (IPUM); horizons and catalogue numbers are given in the figure captions.

Phylum Chordata Bateson, 1886

Class Conodonta Pander, 1856

Order Ozarkodinida Dzik, 1976

Family Spathognathodontodae Hass, 1959

Genus *Ozarkodina* Branson & Mehl, 1933

Type species. – *Ozarkodina confluens* Branson & Mehl, 1933.

Remarks. – The genus *Ozarkodina* was recently subdivided into several new genera by Murphy *et al.* (2004) and Carls *et al.* (2005). We agree with these authors that in the previous meaning *Ozarkodina* included several different groups which may represent separate genera; however, in our opinion, this new scheme may be imperfect as well. *Wurmiella* has been proposed to include taxa related to the "excavata Group", while *Zieglerodina* is the new genus name for the "remscheidensis Group". These subdivisions seem to be valid and the names accepted, apart from some possible nomenclature problems highlighted by Donoghue *et al.* (2008). Murphy *et al.* (2004) also proposed a new genus to represent the "eosteinhornensis Group", which has not been accepted according to ICZN rules. We believe that the apparatus structure and the morphology of the elements in this group are very close to that of the "remscheidensis Group," and therefore could belong to the genus *Zieglerodina*. However, since there is still ongoing debate in the scientific community as to the affinity of this group, and our material is not sufficient to reach a definitive solution, we refer to taxa of the latter group as "*Ozarkodina*" *eosteinhornensis* in the remainder of this paper. A few other taxa, such as "*Ozarkodina*" *snajdri* (Walliser) and "*Ozarkodina*" *crispa* (Walliser) still do not have a home after this subdivision, because their apparatuses remain unknown.

Ozarkodina cf. confluens Branson & Mehl, 1933

Figure 4N

1964 *Spathognathodus primus* (Branson & Mehl). – Walliser, pl. 22, fig. 16 (only).

1975 *Ozarkodina confluens* (Branson & Mehl). – Klapper & Murphy, p. 68, pl. 4, figs 24, 25 (only).

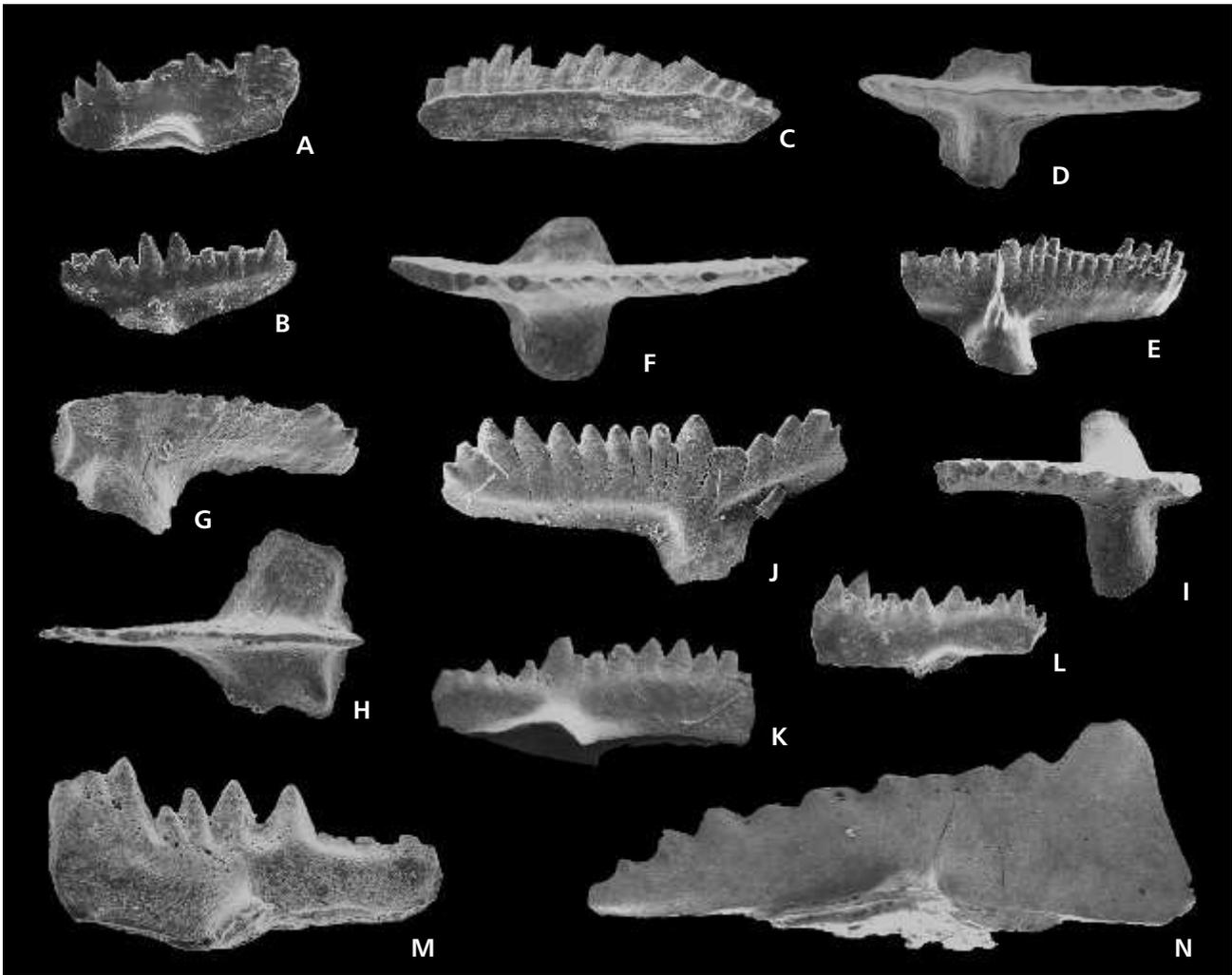


Figure 4. A – *Zieglerodina* cf. *remscheidensis* (Ziegler, 1960); IPUM 28203, lateral view of Pa element, sample MC II 5, *detortus* Zone; × 70. • B – *Zieglerodina* cf. *remscheidensis* (Ziegler, 1960); IPUM 28204, lateral view of Pa element, sample MC II 5 *detortus* Zone, × 70. • C – *Wurmiella excavata* (Branson & Mehl, 1933); IPUM 27671, lateral view of Pa element; sample MC II 1, *Oz. crista* Zone; × 70. • D – *Ozarkodina eosteinhornensis* s.s. (Walliser, 1964); IPUM 27673, upper view of Pa element; sample MC II 3D, *eosteinhornensis* s.l. Zone; × 70. E – *Ozarkodina eosteinhornensis* s.s. (Walliser, 1964); IPUM 27674, lateral views of Pa element; sample MC II 3D, *eosteinhornensis* s.l. Zone; × 70. F – *Ozarkodina eosteinhornensis* s.l. (Walliser, 1964); IPUM 27672, upper view of Pa element; sample MC II 3D, *eosteinhornensis* s.l. Zone; × 70. • G, H – *Ozarkodina crista* (Walliser, 1964); IPUM 27677, lateral (G) and upper (H) views of Pa element; sample MC II 1, *crispata* Zone; × 70. • I – *Lanea* cf. *omoalpha* (Murphy & Valenzuela-Rios, 1999); IPUM 28205 upper view of Pa element, sample MC II 6, *woschmidti* Zone; × 70. • J – *Ozarkodina planilingua* (Murphy & Valenzuela-Rios, 1999); IPUM 28192, upper-lateral view of Pa element; sample MC II 5F, *woschmidti* Zone; × 70. • K – *Ozarkodina planilingua* (Murphy & Valenzuela-Rios, 1999); IPUM 28191, lateral view of Pa element; sample MC II 5F, *woschmidti* Zone; × 70. • L – *Ozarkodina confluens* (Branson & Mehl, 1933); IPUM 27680, lateral view of Pa element; sample MC II 0, *snajdri* Zone; × 70. • M – *Ozarkodina confluens* (Branson & Mehl, 1933); IPUM 27681, lateral view of Pa element; sample MC II 4A, *eosteinhornensis* s.l. Zone; × 70. • N – *Ozarkodina* cf. *confluens* (Branson & Mehl, 1933); IPUM 28193, lateral view of Pa element; sample MC II 4A, *detortus* Zone; × 45.

Description. – The single Pa element recovered has a general triangular outline in lateral view, since the blade is definitely higher in the anterior part and declines steeply towards the posterior end. A distinct set of two fused denticles is located in the anterior part of the blade; nine denticles are on the remaining part of the blade; denticles well spaced, almost fused and laterally compressed. The subtriangular basal cavity is located under the central part of the element.

Remarks. – The studied specimen is close to *Oz. confluens* in the general shape, but the denticles pattern and the strongly subtriangular outline of the blade are not typical; therefore, we prefer to name it *Oz. cf. confluens*. It is similar to representatives of genus *Amydrotaxis* in having spaced and partially fused denticles, but differs in the shape of basal cavity, which does not reach the extremities of the blade. Similar elements have been figured by Walliser (1964), and reported as *Oz. confluens* delta morphotype by Klapper & Murphy (1975).

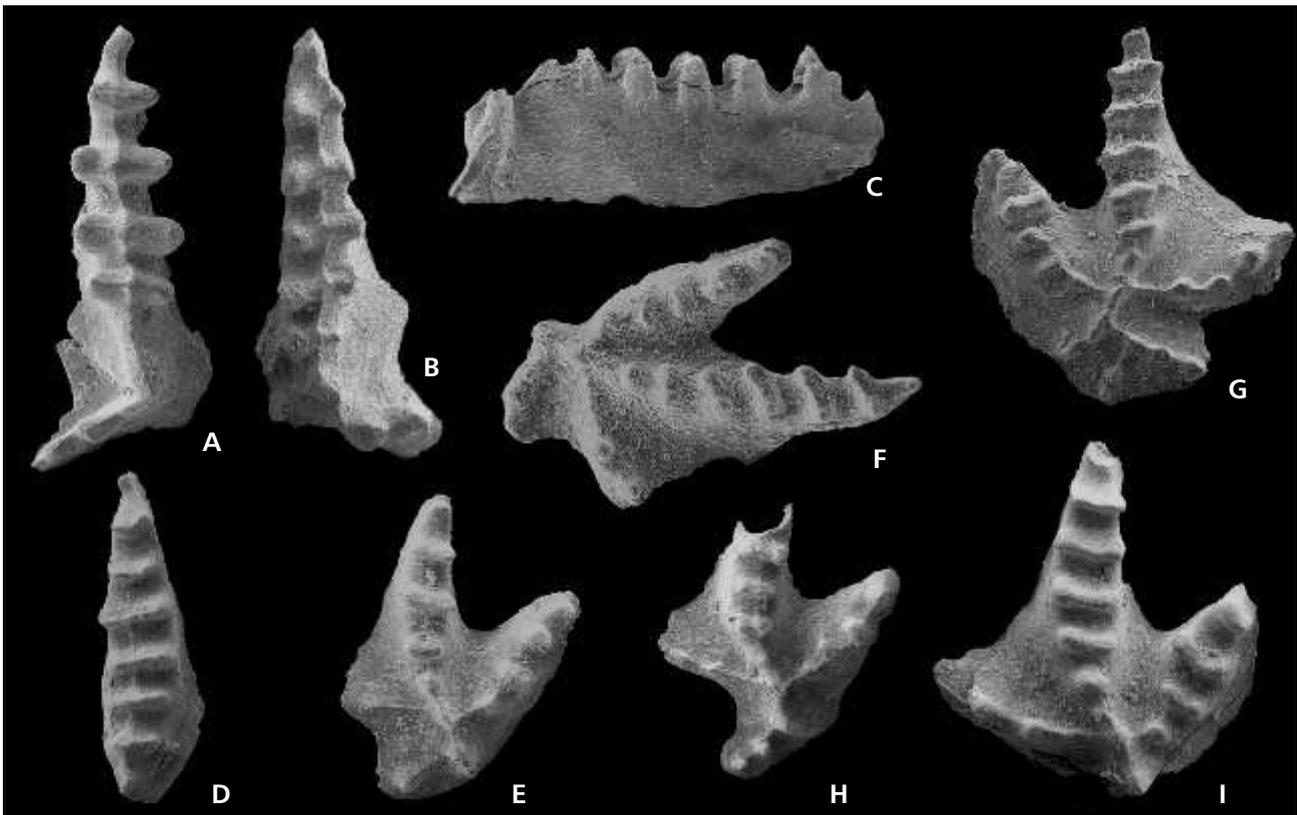


Figure 5. A – *Icriodus waschmidti waschmidti* (Ziegler, 1960); IPUM 28196, upper view; sample MC II 5F, *woschmidti* Zone; × 70. • B, C – *Icriodus waschmidti hesperius* (Klapper & Murphy, 1975); IPUM 28195, upper (B) and lateral (C) views; sample MC II 5F, *woschmidti* Zone; × 45. • D – *Icriodus* sp.; IPUM 28197, upper view; sample MC II 6, *woschmidti* Zone; × 70. • E – *Pedavis biexoramus* (Murphy & Matti, 1982); IPUM 28201, upper view; sample MC II 5F, *woschmidti* Zone; × 70. • F – *Pedavis biexoramus* (Murphy & Matti, 1982); IPUM 28200, upper view; sample MC II 5F, *woschmidti* Zone; × 70. • G – *Pedavis biexoramus* (Murphy & Matti, 1982); IPUM 28199, upper view; sample MC II 5F, *woschmidti* Zone; × 70. • H – *Pedavis biexoramus* (Murphy & Matti, 1982); IPUM 28202, upper view; sample MC II 6, *woschmidti* Zone; × 70. • I – *Pedavis biexoramus* (Murphy & Matti, 1982); IPUM 28198, upper view; sample MC II 6, *woschmidti* Zone; × 70.

Stratigraphic range. – The studied specimen comes from the *eosteinhornensis* s.l. interval Zone. The specimen figured by Walliser (1964), also from the Carnic Alps, is from the same stratigraphic level, whereas specimens figured by Klapper & Murphy (1975) are definitely older.

Material. – 1 Pa element from sample MC II 4A.

Order Prioniodontida Dzik, 1976
Family Icriodontidae Müller & Müller, 1957

Genus *Pedavis* Klapper & Philip, 1971

Type species. – *Icriodus pesavis* Bischoff & Sannemann, 1958.

***Pedavis biexoramus* Murphy & Matti, 1982**
Figures 5E–I

1975 *Pedavis* sp. nov. C. – Klapper & Murphy, p. 50, pl. 12, fig. 12.

1979 *Pedavis* sp. – Lane & Ormiston, pl. 1, fig. 14.

1980 *Pedavis* cf. *biexoramus* Murphy *et al.* – Schönlaub, pl. 1, fig. 14.

1982 *Pedavis biexoramus* n. sp. – Murphy & Matti, pp. 51–54, pl. 5, figs 28, 33.

1992 *Pedavis biexoramus* Murphy & Matti. – Barrick & Klapper, p. 45, pl. 3, figs 1–3.

Remarks. – The specimens of *P. biexoramus* from the Monte Cocco II section are similar to the type material described by Murphy & Matti (1982). All have well developed inner and outer processes, except the specimen in Fig. 6G, which has a weak process.

Stratigraphic range. – *P. biexoramus* is known from the lowermost Devonian of the Carnic Alps from the Oberbuchach 2 Section [slightly above the last occurrence of *Icriodus waschmidti* (Schönlaub, 1980)], and in Nevada from within the *eurekaensis* Zone (Murphy & Matti 1982). However, a possible earlier origin of the species within the *hesperius* Zone is suggested by Murphy & Matti (1982, fig. 9),

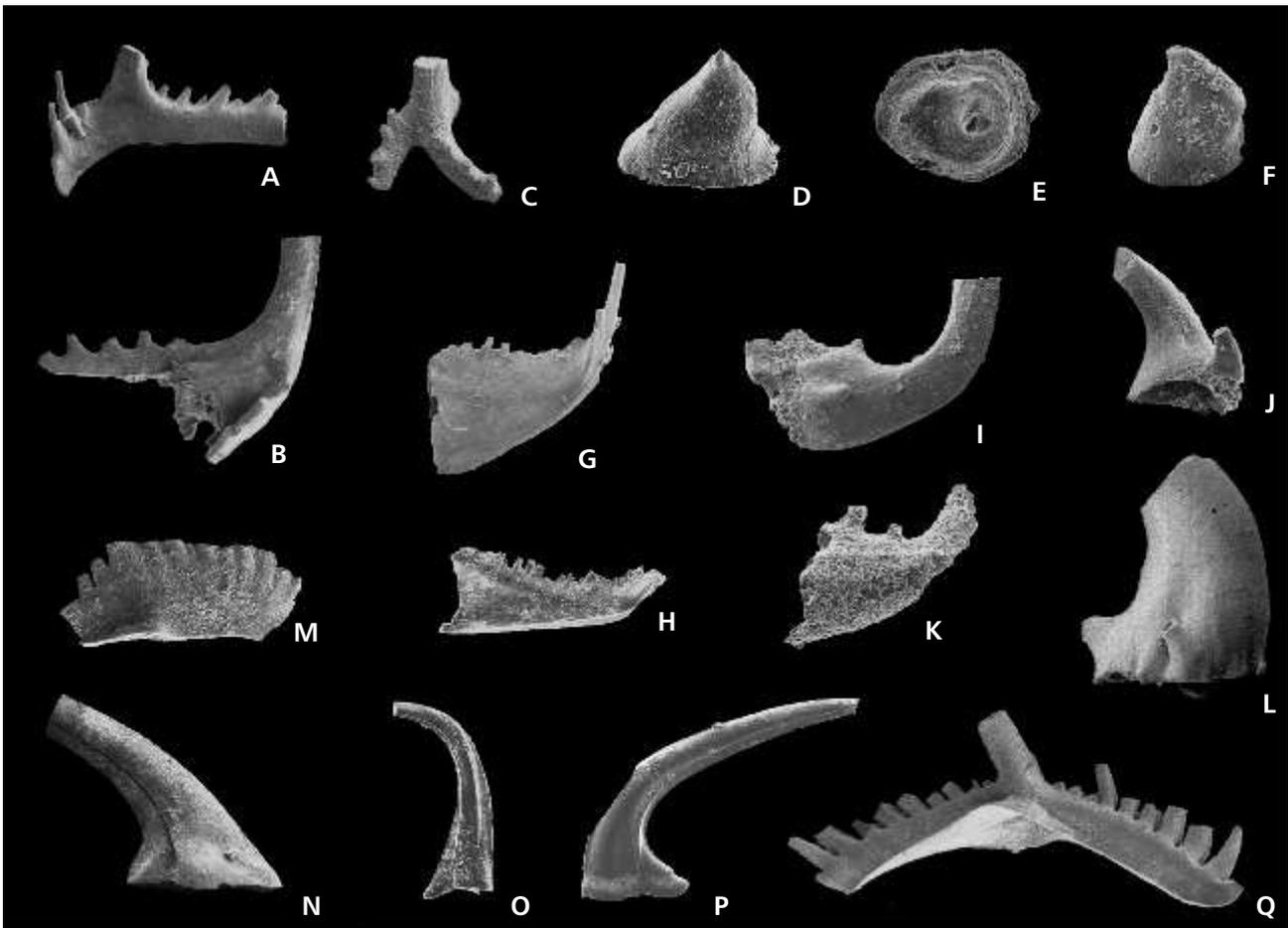


Figure 6. A – *Oulodus elegans detortus* (Walliser, 1964); IPUM 28180, lateral view of Sc element, sample MC II 3D, *detortus* Zone; × 70. • B – *Oulodus elegans elegans* (Walliser, 1964); IPUM 28181, lateral view of Sc element, sample MC II 5, *detortus* Zone; × 70. • C – *Oulodus elegans elegans* (Walliser, 1964); IPUM 28182, posterior view of Pb element, sample MC II 3C, *detortus* Zone; × 70. • D, E – *Pseudooneotodus beckmanni* (Bischoff & Sannemann, 1958); IPUM 27678, lateral (D) and upper (E) views, sample MC II 1, *crispa* Zone; × 90. • F – *Pseudooneotodus bicornis contiguus* (Corradini, 2008); IPUM 28183, lateral view; sample MC II 1B, *eosteinhornensis* s.l. Zone; × 90. • G – *Belodella anomalis* (Cooper, 1974); IPUM 28184, lateral view; sample MC II 3D, *detortus* Zone; × 70. • H – *Belodella resima* (Philip, 1965); IPUM 27675, lateral view of Pa element; sample MC II 4, *eosteinhornensis* s.l. Zone; × 70. • I – *Coryssognathus dubius* (Rhodes, 1953); IPUM 28187, lateral view of Sa/Sb element; sample MC II 4, *detortus* Zone; × 70. • J – *Coryssognathus dubius* (Rhodes, 1953); IPUM 28186, lateral view of M element; sample MC II 4, *detortus* Zone; × 70. • K – *Coryssognathus dubius* (Rhodes, 1953); IPUM 27683, lateral view of Sc element; sample MC II 4, *detortus* Zone; × 70. • L – *Coryssognathus dubius* (Rhodes, 1953); IPUM 28188, lateral view of Sa/Sb element; sample MC II 4, *detortus* Zone; × 70. • M – *Pelekysgnathus* sp. A (Thomas, 1949); IPUM 28190, lateral view of Pa element; sample MC II 1, *crispa* Zone; × 70. • N – *Dapsilodus obliquicostatus* (Branson & Mehl, 1933); IPUM 28189, lateral view; sample MC II 0, *crispa* Zone; × 50. • O – *Panderodus unicastatus* (Branson & Mehl, 1933); IPUM 28185, lateral view; sample MC II 1, *crispa* Zone; × 50. • P – *Panderodus recurvatus* (Branson & Mehl, 1933); IPUM 28206, lateral view; sample MC II 1, *crispa* Zone; × 50. • Q – *Oulodus siluricus* (Branson & Mehl, 1933); IPUM 28194, lateral view of Pb element; sample MC II 5C, *detortus* Zone; × 45.

and also on the basis of joint occurrences at Copenhagen Canyon V-IV (Murphy & Matti 1982, tab. 1). Our material appears to confirm their hypothesis, because it comes from samples where *Icriodus hesperius* is also present.

Material. – 6 Pa elements from samples MC II 5F-MCII 5.

Genus *Pelekysgnathus* Thomas, 1949

Type species. – *Pelekysgnathus inclinatus* Thomas, 1949.

Pelekysgnathus sp. A

Figure 6M

1995 *Pelekysgnathus index* Klapper & Murphy. – Barca *et al.*, pl. 4, figs 4, 5.

1998 *Pelekysgnathus index* Klapper & Murphy. – Serpagli *et al.*, p. 110, pl. 1.2.2, figs 4, 5

Description. – The element is small, laterally compressed and with the anterior end slightly bent downward. The blade bears denticles that are partially fused, laterally com-

pressed and of different size. The posteriormost denticle is strongly posteriorly directed and looks bigger than the other denticles, even if broken. The basal cavity is symmetrical, oval and not very wide; it is located under the posterior third of the element.

Remarks. – The posterior position of the basal cavity and the features of the posterior denticle allow us to refer the specimen to the Genus *Pelekysgnathus*. This taxon differs from *P. index* in having a smaller and almost symmetrical basal cavity, a less developed posterior denticle and much more differentiated denticles.

Similar specimens have been found in Sardinia by Barca *et al.* (1995) and Serpagli *et al.* (1998) from the lowermost part of the *crispa* Zone.

Stratigraphic range. – Our specimen and those figured in literature are all from the lower part of the *crispa* Zone.

Material. – 1 Pa element from samples MC II 1.

Family Pterospodontidae Cooper, 1977

Genus *Coryssognathus* Link & Druce, 1972

Type species. – *Cordylodus? dubius* Rhodes, 1953.

Coryssognathus dubius (Rhodes, 1953)

Figures 6I–L

- 1993 *Coryssognathus dubius* (Rhodes). – Miller & Aldridge, p. 243–253, pl. 1, figs 1–14; pl. 2, figs 1–18; pl. 3, figs 1–12; pl. 4, figs 1–15; text-figs 2–6 (*cum syn.*).
- 1997 *Coryssognathus dubius* (Rhodes). – Serpagli *et al.*, pp. 240–242, pl. 1, figs 1–15 (*cum syn.*).
- 1998 *Coryssognathus dubius* (Rhodes). – Corradini *et al.*, pl. 3.3.1, fig. 15.
- 1999 *Coryssognathus dubius* (Rhodes). – Viira, pl. 1, figs 9–12.
- 1999 *Coryssognathus dubius* (Rhodes). – Cockle, p. 120, pl. 4, fig. 20.
- 2001 *Coryssognathus dubius* (Rhodes). – Corradini *et al.*, pl. 1, figs 10, 11.
- 2003 *Coryssognathus dubius* (Rhodes). – Corradini *et al.*, pl. 1, fig. 13.
- 2004 *Coryssognathus dubius* (Rhodes). – Farrell, pp. 959, 960, pl. 4, figs 1–9.

Remarks. – The apparatus of *Coryssognathus dubius* was reconstructed as septimembrate by Miller & Aldridge (1993) and their proposal was followed by Serpagli *et al.* (1997). The taxon is documented in several palaeoconti-

nents, but only from Llandovery-Ludlow; however, in northern Gondwana it appears to have a longer range, nearly reaching the top of the Silurian (*i.e.* Sardinia, Serpagli *et al.* 1997). The taxon occurs in the Monte Cocco II section up to the lower part of the *detortus* Zone, and is found in levels where *Oz. eosteinhornensis* s.s. has also been recovered, confirming the range documented in Sardinia. Therefore, it appears that *C. dubius* has a longer range in northern Gondwana than elsewhere.

Stratigraphic range. – The species is known from the late Llandovery (Mabillard & Aldridge 1983) to the uppermost Silurian (Serpagli *et al.* 1997, and this paper).

Material. – 1 Pc, 5 M, 6 Sa/Sb, 3 Sc and 5 coniform elements, from samples MC II 2, MC II 3, MC II 4, MC II 4A and MC II 5.

Conclusions

The main results of this paper can be summarized as follows:

1. the conodont association allows the discrimination of five late Silurian-lowermost Devonian conodont zones: *snajdri*, *crispa*, *eosteinhornensis* s.l., *detortus* and *woschmidtii*;

2. the Silurian/Devonian boundary occurs in the upper part of the section. However, it is not possible to locate it accurately; it can be constrained within a small interval between the last occurrence of *Oz. confluens* (MC II 5D) and the first occurrence of Icriodontids (MC II 6). Furthermore, the occurrence of typical forms of *Z. remscheidensis* in sample MC II 5E indicates that this level is likely Devonian in age;

3. a new species of *Pelekysgnathus* is described from the *crispa* Zone, and left in open nomenclature. This taxon is potentially useful for regional correlation within northern Gondwana;

4. *Coryssognathus dubius* reaches the upper part of the *detortus* Zone, occurring together with *Oz. eosteinhornensis* s.s., confirming its longer range (up to the Přídolí) in northern Gondwana.

Acknowledgements

Luca Simonetto (Udine), Monica Pondrelli (Pescara) and Paolo Serventi (Modena) shared the work in the field, providing useful comments on the geology of the area. Angelo Ibba (Cagliari) and

Claudio Gentilini (Modena) took scanning microscope pictures and Salvatore Vacca (Cagliari) helped in the laboratory. Mike Murphy (Riverside) and Ladislav Slavík (Prague) critically read the manuscript providing useful comments. This paper is a contribution to IGCP Project No. 499 “Devonian sea-land interactions: evolution of ecosystem and climate”.

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